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Rabett, R. (2017). Tropical limestone forest resilience and late Pleistocene foraging during MIS-2 in the Trảng An massif, Vietnam. *Quaternary International*, 448, 62-81. <https://doi.org/10.1016/j.quaint.2016.06.010>

**Published in:**  
Quaternary International

**Document Version:**  
Peer reviewed version

**Queen's University Belfast - Research Portal:**  
[Link to publication record in Queen's University Belfast Research Portal](#)

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# TROPICAL LIMESTONE FOREST RESILIENCE AND LATE PLEISTOCENE FORAGING DURING MIS-2 IN THE TRÀNG AN MASSIF, VIETNAM

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## ABSTRACT

*In this paper we present a multi-proxy study of tropical limestone forest and its utilization by human groups during the major climatic and environmental upheavals of MIS-2 (29-11.7 kBP). Our data are drawn from new field research within the Trảng An World Heritage property on the edge of the Red River Delta, northern Vietnam. Key findings from this study include <sup>1)</sup> that limestone forest formations were resilient to the large-scale landscape transformation of the Sunda continent at the end of the last glaciation; <sup>2)</sup> that prehistoric human groups were probably present in this habitat through-out MIS-2; and <sup>3)</sup> that the forested, insular, karst of Trảng An provided foragers with a stable resource-base in a wider changing landscape. These results have implications for our understanding of the prehistoric utilization of karst environments, and resonance for their conservation in the face of climate and environmental change today.*

Keywords: limestone forest, resilience, MIS-2, tropical foragers, refugium, conservation

*Quaternary International* (in press) <http://dx.doi.org/10.1016/j.quaint.2016.06.010>

## 1. INTRODUCTION

Profound climate and landscape changes were witnessed at a global scale during Marine Isotope Stage 2 (MIS-2) 29-11.7 cal. kBP. As conditions deteriorated into the Last Glacial Maximum (LGM) 26-19 cal. kBP (Clark *et al.* 2009) heightened aridity was experienced in many parts of the dry tropics (e.g. Weyhenmeyer *et al.* 2000; Barton *et al.* 2005; Cook 2009). Both here and at higher latitudes (that were more directly under the influence of LGM ice-sheet expansion) biotic and human communities tended to contract into refugial enclaves (see e.g. Pala *et al.* 2012; Rose *et al.* 2013; Williams *et al.* 2013; Husemann *et al.* 2014).

While the effects of the LGM were more muted in the humid tropics of Southeast Asia (e.g. Huang *et al.* 1997; Pelejero *et al.* 1999; Visser *et al.* 2003; Shintani *et al.* 2011) local marine, terrestrial and lacustrine archives indicate that changes in regional precipitation and vegetation cover did, nonetheless, occur (Huang *et al.* 1997; Sun & Li 1999; Sun *et al.* 2000). In the post-LGM late glacial especially, climatic variability and environmental changes had far-reaching and transformative effects on Pleistocene landscapes. The increased aridity associated with Heinrich Event 1 (H1) (c. 17.5-14.7 cal. kBP) brought tangible climatic impact (Wang *et al.* 2001; Huang *et al.* 2011; Marwick & Gagan 2011; Standford *et al.* 2011; Partin *et al.* 2015). The final resurgence of near-fully glacial conditions during GS-1 (Younger Dryas) 12.7-11.7 cal. kBP in the circum-Atlantic does not seem to have impacted on equatorial latitudes in Southeast Asia (Partin *et al.* 2007), though its effect at more northerly tropical latitudes is tenable (Huang *et al.* 2011).

The process of deglaciation also initiated sustained sea-level rise and flooding of regional coasts. With rates of inundation ranging from c. 0.4 m to as much as 5.0 m per 100 years, the effect on human and biological communities would have been considerable in Southeast Asia (Hanebuth *et al.* 2000; 2009; Kienast *et al.* 2003; Schimanski & Stattegger 2005; Verleyehn *et al.* 2005; Soares *et al.* 2008; Kopp 2012; Hunt & Gilbertson 2014). Ultimately, c. 75 per cent of the region's low-lying Pleistocene landmass, 'Sunda Land' (Molengraaf & Weber 1921) – incorporating the Greater Sunda Islands (Borneo, Sumatra and Java) together with numerous smaller islands and island chains, which extended Asia into the Southern Hemisphere – was submerged by the mid-Holocene. Across this continent and latter-day archipelago, limestone karst environments have played a prominent role in early human and archaic hominin occupation for almost a hundred millennia (e.g. Brown *et al.* 2004; Yi *et al.* 2008; van den Bergh *et al.* 2009; Mijares *et al.* 2010; Rabett 2012; Stimpson 2012; Barker 2013).

Investigation of karstic habitats in northern Vietnam has been closely tied to one particular palaeo-cultural entity: the 'Hoabinhian'. Various described as, a Mesolithic culture (e.g. Matthews 1966), a 'technocomplex' (e.g. Gorman 1970; Forestier *et al.* 2015) reflecting a common ecological adaptation, or an industry (see Shoocongdej 1996), it is now argued that the Hoabinhian represents the most important point of cultural reference for the late glacial occupation of the seasonal tropics (Ji *et al.* in press). It is certainly the case that it has featured conspicuously in our understanding of early human adaptation to these environments through almost a century of research (Gorman 1970, 1971; Reynolds 1989; Ha Van Tan 1997; Shoocongdej 1996; Rabett 2012). Despite this attention, the significant technological and subsistence variability that exists within the Hoabinhian continues to be a point of debate and reinterpretation (e.g. Rabett *et al.* 2011; Marwick 2013; Forestier *et al.* 2013, 2015). Although generally considered to have emerged as a distinct technological system during, or soon after, the LGM, lasting into the Holocene, the duration of the Hoabinhian and its geographic extent (e.g. Shoocongdej 1996; Ji *et al.* in press) are also yet to

be formally agreed upon. Its heartland, though, continues to be situated today, much as it was following initial discoveries by Patte in 1923 and Colani and Mansuy (1926-31): in the forest-clad karstic hills of northern Vietnam; and its emergence coming through the process of living in that environment.

Tropical limestone forests differ significantly in structure and species composition from other tropical forest formations (Sterling *et al.* 2006). The insular and refugial character they exhibit reflects high levels of biodiversity and species endemism (e.g. Adam & Mamat 2005; Clements *et al.* 2006, 2008; Sodhi *et al.* 2007; Furey *et al.* 2010; Gao *et al.* 2015). Soils are generally thin, poor in most minerals and nutrients, and derived directly from the underlying limestone. Modern precipitation levels in the northern sub-tropics, close to the Song Hong River delta of Vietnam, are concentrated (c. 80%) during the southwest monsoon between May and October, where receipt of rainfall ranges from 1300-1800 mm (Li *et al.* 2006b). The colder and drier northeast monsoon between November and April brings much lower levels of annual rainfall: 200-250 mm (Do Cong Thung 2014). Surface water run-off is often rapid in karst landscapes. This has the effect of heightening local aridity. Surface water accumulation in poorly drained locales, though, can lead to the formation of acidic peat, leading to the creation of varied microhabitats and the ability of limestone areas to host vegetative components adapted to a range of different conditions (Sterling *et al.* 2006). While precipitation arrives with marked monsoon seasonality and is often poorly retained for plant-life at sub-tropical latitudes, compared to their southern lowland rainforest counterparts (Chang *et al.*, 2005) summer insolation levels are routinely more intense than those experienced closer to the Equator (Paillard *et al.* 1996; Wang & Enfield 2001). This combination has a profound effect on the production of plant *n*-alkanes and *n*-alkanoic acids that can be tracked and through them past environmental conditions can be reconstructed (Rommerskirchen *et al.* 2003).

The particular association of northern Vietnam's palaeo-cultural record with tropical limestone forest emphasises that greater consideration needs to be made of the diversity that exists in tropical forest ecologies and the impact this has had on the history and scope of human adaptation (e.g. Roberts & Petraglia 2015). With recent claims that less is currently known about human adaptive capacities than about the responses to climate change by many natural systems (IPCC 2007-II), examining the long-term relationship between humans and tropical forest habitats offers considerable potential to enhance our understanding of the diversity of socio-cultural responses to climate-driven environmental change. In this paper we present a multi-proxy terrestrial record from the cave site of Hang Trống, located in an isolated limestone massif on the western edge of the Song Hong (Red River) delta, Vietnam. This record spans the last twelve thousand years of MIS-2 and affords a rare opportunity to assess the local impact of late glacial climate and environmental change on northern tropical limestone forest; also to consider the significance of this habitat to early human settlement of the area; and to explore the potential value of such Pleistocene records to modern conservation management. The Trảng An limestone massif, where this study is set, is recognized internationally as an exemplar of the combined natural and cultural heritage value preserved within such landscapes.

## 2. TRẢNG AN LANDSCAPE COMPLEX

This World Heritage property is situated within the province of Ninh Bình in northern Vietnam, on the southern margin of the Song Hong (Red River) Delta, and is centred on an isolated massif of Upper Palaeozoic massively-bedded limestone karst of shallow marine

lagoon and reef origin (figure 1). The core property is forested and covers 6226 hectares or 62.2 km<sup>2</sup> (surrounded by a buffer zone of 6026 hectares, of mostly rural land with rice paddy fields). Trảng An is administered by the Ninh Binh Province People's Committee and a semi-autonomous management board. The property exemplifies an outstanding humid tropical tower-karst landscape in the final stages of geomorphic evolution, comprising karst cones and towers, enclosed depressions and an intricate system of fresh water-filled foot-caves. Although currently emergent, this landscape has been invaded by the sea several times in the recent geological past and is often compared to the flooded tower karst seascape of Halong Bay. Trảng An's inscription also recognizes its prehistoric archaeology, which spans in excess of the last 30,000 years and contains one of the country's oldest archaeological sites: Hang Trống.

<< Figure 1 >>

### 3. HANG TRỐNG

The cave site of Hang Trống (20.250444N, 105.890111E) is situated approximately 142.3 metres above sea level, and 60 m below the apex of a limestone karst tower in the central part of the Trảng An massif. The cave forms a tunnel that is orientated along a north-south axis through the apex, with a total exposed floor area of *c.* 170 m<sup>2</sup> (figure 2). Hang Trống's elevated position, coupled with a marked difference in the size of its two apertures, promotes air flow, accelerating it substantially when winds are northerly. Aeolian processes have been a significant force in site formation: affecting not only rates of deposition but also rates of the erosion and transportation of sediments from the cave (probably exiting via the southern aperture). The noticeable loss of *c.* 1 m of sediment (indicated by surviving material adhering to the walls), was probably lost through a mix of hydrological and aeolian erosion.

<< Figure 2 >>

#### 3.1 Chronology

Radiocarbon dates have been obtained using the accelerator mass spectrometer (AMS) at the 14CHRONO Centre, Queen's University, Belfast (table 1 – <sup>14</sup>C ages were calibrated by Calib 6.0 on the 2<sup>nd</sup> January 2013, using the INTCAL09 calibration curve, [Reimer \*et al.\* 2009](#)). All ages are quoted in calibrated age before present (BP), with BP being AD 1<sup>st</sup> January 1950; the abbreviation 'kBP' is used to describe calibrated thousands of years. Aside from one surface-modern age from mixed deposits, all <sup>14</sup>C dates are in sequence (see figure 3) and demonstrate the potential for accurately determining a time-depth profile at the site, pushing evidence of human occupation here deep into the LGM. The aeolian processes at work within Hang Trống have probably aided the chronological stratigraphy, bringing sediments in to Hang Trống regularly and preventing post-depositional mixing of the shell midden. The onset of the LGM coincides with the base of excavation (2010) in Trench 1.

<< Table 1 >>

### 4. METHODOLOGIES AND RESULTS

Hang Trống was surveyed using a Geo Fennell 20x automatic level and mapped at high resolution using a Leica HDS3000 LiDAR scanner. Systematic excavation began in the



middle of the cave with Trench 1 in 2009 and then, in the second season (2010), concurrently in two additional trenches, Trenches 2 and 3, close to the western and eastern walls, respectively (figure 3) (see [Rabett 2013](#)). All three trenches contained cultural evidence, including hearths. In addition to macroscopic data from these material remains, geochemical analysis was undertaken on the sediments (and shell), providing a multi proxy approach to interpreting the location, its occupation and local environment.

### << Figure 3 >>

#### 4.1 Macro-botanical remains

Samples of macro-botanical remains were collected from dry sieving (2 mm mesh) on site and also in bulk sediment samples that had been processed using bucket flotation (1 mm). Data on carbonized plant remains, in particular, provides a useful line of evidence on vegetation composition, fuel collection and plant processing strategies ([Ceron n.d.](#)). In each case samples were counted, weighed, measured and assessed for state of preservation. A total of 459 fragments were identified to family or higher taxonomic level (table 2). Specimens were mounted and photographed with a NIKON Eclipse Ci with 4x.15 to 100xA/1.25 magnification. The taxonomic identifications were based on comparisons with identified specimens, from the nearby Vietnamese Palaeolithic site of Con Moong ([Nguyen Viet pers. comm. to Ceron 2011](#)), from Ille Cave, Palawan ([Carlos 2010](#); [Ceron pers. observ.](#)), on published material ([Paz 2001](#)), and with the assistance of the Forest Product Research and Development Institute (FPRDI), Los Baños, Laguna. Following established criteria relating to characteristics such as pore topography, tracheids and fibre (e.g. [Thompson 1996](#); [Ella et al. 2009](#)) a pilot SEM study of four pieces of charred parenchymous material was also undertaken. The chosen material had been recovered through flotation from contexts: (8002), (8101), (8103) from Trench 1 at Hang Trống and from (8215) from Trench 2.

### << Table 2 macro-botanical >>

#### 4.2 Palynology

Sediment samples were collected from the north-facing section of Trench 1 for palynological analysis. The surface of the section was cut back to remove loose material and ensure there was no modern contamination. Palynological analysis has been widely used as a method to determine past vegetation history (e.g. [Birks 1980](#); [Birks & Birks 1981](#)). In upland cave sites the sediments are dry and highly oxic; unlike in water lain environments where available oxygen (and hence degradation) is limited. As a result pollen grains oxidize rapidly and/or become 'crumpled'. This level of degradation of pollen grains can pose problems to the reliability when standard palynological techniques (those applied here) are employed. Due to the high concentration of cyclophorids at the site, all samples were sieved to remove shell fragments prior to standard pollen analysis (after [Fægri & Iversen, 1989](#)). Pollen was successfully found in all six of the samples investigated, though in varying concentrations as shown in table 3. Owing to the importance of aeolian processes in site formation here, the pollen record is likely to be drawn from a wide catchment, extending beyond Trảng An onto the plains that surround the massif.

Despite the degraded state of pollen at the site, unique identifying features could still be recognized (table 4). The top context sample (8000-8002) displays a considerably higher concentration of identifiable pollen than other contexts (a value of 60 % recoverable

identification). It is deduced that this better preservation is largely due to modern wind-blown pollen becoming trapped within the sediments as erosion and localised micro-re-deposition of these takes place. Consequently, these data could provide an example of the modern vegetation surrounding the site, though admixture with Pleistocene sediments has probably distorted this picture. The underlying and, by comparison, undisturbed Pleistocene deposits exhibited a much reduced level of preservation, with only 2-6 % of recovered pollen being identifiable using standard preparation techniques. Nonetheless, a range of plant types was identified and provided a workable dataset.

<< Table 3 pollen counts >>

<< Table 4 pollen identification >>

#### 4.3 Plant biogeochemical markers

Cave locations can provide chemically stable environments suitable for the preservation of biogeochemical compounds from higher plants (Blyth *et al.* 2007; Huang *et al.* 2008). Plant biomarkers are released into the environment by plant metabolic processes and decomposition; they can be isolated from other organic molecules and used as environmental markers (see Eglinton & Hamilton 1967; Van Bergen *et al.* 1997). *n*-Alkanes in plants are hydrocarbons that are mainly synthesized in the leaf wax for protection against incoming solar radiation (insolation), desiccation and pathogens (see e.g. Hadley 1980; Koch & Ensikat 2008; Kosma *et al.* 2009). The longer the *n*-alkane chain length the more energy and resources are required for creation, but the greater the protection provided (Rommerskirchen *et al.* 2003). Within equatorial regions where insolation is high plants produce the long *n*-alkane chains, provided there is also sufficient moisture. During periods of climate change that place stress on plants, they either produce shorter *n*-alkane chain lengths or are unable to survive. *n*-Alkanes preserve well in Quaternary deposits, and as the environment and climate are important controls on chain length distribution, they can furnish an independent and complementary record into past conditions (Meyers & Ishiwatari 1993; Schwark *et al.* 2002; Blyth *et al.* 2007; Ludgate 2013).

Bulk and compound-specific  $\delta^{13}\text{C}$  analyses of sedimentary organic carbon preserved *n*-alkanes and *n*-alkanoic acids and provide information about the photosynthetic pathway of the plants they derive from and, therefore, the environment in which the plant grew. Ranges in carbon isotope values represent plants using different photosynthetic pathways. The  $\delta^{13}\text{C}$  value can range from *c.* -35 to *c.* -15 ‰ (relative to the Vienna Pee Dee Belemnite international standard) depending on a number of variables, including (but not limited to) plant genus, moisture stress or plant longevity (Quade *et al.* 1989; Lockheart *et al.* 1997). Plants using the  $\text{C}_3$  photosynthetic pathway require sufficient amounts of water during their growing season to photosynthesise and give  $\delta^{13}\text{C}$  values typically within the range -25 ‰ to -35 ‰ (O'Leary 1981; Osborne & Slack 2012). Plants that follow the  $\text{C}_4$  photosynthetic pathway (including, in particular, most grass families) are able to survive in more arid conditions.  $\text{C}_4$  plants have an average  $\delta^{13}\text{C}$  value of -13 ‰ and range between -12 ‰ to -16 ‰ (O'Leary 1981; Osborne & Slack 2012).

Rabett *et al.* (2011) demonstrate that plant biomarkers, preserved for over 10 kBP in sediments at another cave site in Trảng An, Hang Boi (*c.* 1 km from Hang Trống), display similarity to the biomarkers of plants within the surrounding vegetation today. For the biogeochemical analysis at Hang Trống samples were collected from the North face of

Trench 1. The section surface was cleaned to remove cross-context contamination and some whole shells were removed to increase the mass of sediment. In the laboratory the bulk lipid fraction was soxhlet-extracted from the sediment using dichloromethane (95%) and methanol (5%) to maximise recovery (Banjoo & Nelson 2005; Wang *et al.* 2010). The recovered organic matrix was then analysed using Agilent 6890 Series gas chromatogram coupled to an Agilent 5973 mass spectrometer (GC–MS). Chemstation software was utilised to identify and quantify the *n*-alkanes.  $\delta^{13}\text{C}$  values of *n*-Alkanoic acids was determined using Delta V Advantage ThermoFisher isotope-ratio mass spectrometer interfaced with GC-Isolink Trace GC Combustion conversion system. *n*-Alkanes were removed using off-line chromatography and analysed for compound specific  $\delta^{13}\text{C}$  using Agilent gas chromatogram coupled to a Thermo MAT 253 isotope ratio mass spectrometer (GC–IR-MS). Sediment total organic carbon content and total organic carbon  $\delta^{13}\text{C}$  values were determined using a Thermo Flash High Temperature Elemental Analyzer (Flash EA) coupled to a Thermo MAT 253 isotope ratio mass spectrometer (IR-MS). All  $\delta^{13}\text{C}$  values were normalized to the Vienna Pee Dee Belemnite standard.

*n*-Alkanes recovered from sediments have a greater proportion of odd-*n*-alkane carbon chain lengths rather than even; indicating that the sediments at Hang Trống contain a robust plant-derived lipid signal, similar to that observed at Hang Boi. The most abundant *n*-alkanes (over 10% of the total measured) include  $\text{C}_{29}$ ,  $\text{C}_{31}$  and  $\text{C}_{33}$ , these are all known to be of higher plant origin and therefore of importance here. *n*-Alkane  $\text{C}_{31}$  is the most dominant through the sediment column, with an average value of *c.* 40 per cent. This demonstrates that plants growing locally are subjected to some environmental stress (high levels of insolation) but, generally, the nutrients and water required for the growth of protective lipids occur in abundance. Percentage histograms of the samples (figure 4) demonstrate a consistent trend.

The  $\delta^{13}\text{C}$  value of *n*-alkanes and *n*-alkanoic acids recovered from cave sediments clearly indicate a  $\text{C}_3$  photosynthetic pathway and therefore plants that require comparatively moisture-rich conditions in which to thrive. The slightly higher  $\delta^{13}\text{C}$  value for bulk organic carbon within the sediments highlights that preserved organic material is likely to be derived from multiple sources, including microbial and anthropogenic (Kohn and Cerling 2002; Nguyen Tu 2011). Food sources introduced by early human groups into the cave from farther afield could, for example, have contributed a biasing factor in the recovery plant biomarkers – a vector for which we are seeking further clarification. Our current hypothesis, based on the available evidence, suggests that limestone forest similar to that seen around Hang Trống today has existed in the vicinity of the cave through-out MIS-2.

<< Figure 4 >>

#### 4.4 Zooarchaeological evidence

##### 4.4.1 Macro-vertebrate fauna

Bones from larger vertebrates were a relatively infrequent component of the excavated material in all three trenches (table 5). Analysis of these assemblages showed them to be predominantly mammalian, highly fragmented, desiccated and in many cases coated with sediment adhered by calcium carbonate deposition. All recovered fragments >5 mm in length were counted and wherever possible refitted. A total of 1031 fragments (hereafter, Number of Specimens: 'NSP') were recorded, of which a total of 85 (8.2 %) was found to be



diagnostic (to element and taxa, hereafter, Number of Identified Specimens: 'NISP'). The majority was only identifiable to family, but in exceptional cases higher order identification were possible. Comparative materials included those held within the Grahame Clark laboratory, Cambridge, supplemented by collections of digital images and relevant textual sources (notably, [Lekagul & McNeely 1988](#)). Despite the relatively low occurrence of bones, inspection of the data indicated that burnt fragments were present throughout the excavated sequence (figure 5 & table 5). Both the high degree of fragmentation and the evidence of burning give us confidence in the anthropogenic origin of this material.

#### << Table 5 vertebrate fauna >>

#### << Figure 5 composite section T1 >>

##### 4.4.2 Micro-vertebrate fauna

On-site observations and the characteristics of ground surface concentrations of micro-vertebrate bone near the western wall of the cave strongly suggested that these assemblages accumulated via pellet deposition beneath a contemporary owl roosting site. This record provides a valuable taphonomic control for the range and density of micro-faunal remains recovered during on-site excavations, and how they are to be interpreted.

A total of 2517 bones were identified to element and order from a collected sample of the surface assemblages. Frog bones (Order: Anura) dominated ( $n = 2230/2517$ , or 89%), with all identifiable ilia attributed to the genus *Rana*. All parts of the skeleton were represented and the bones exhibited minimal breakage and weathering, with no signs of burning or any marked damage from digestive acids (after [Andrews 1990](#); [Stahl 1996](#)). The remainder of the sample comprised a minor component of small mammal bones, including *Rattus* sp., *Mus* sp., *Suncus* sp., and *Eonycteris* sp. as well as, as yet, unidentified fish remains. The taxonomic composition of the assemblage with its strong bias toward amphibians, combined with the relative size of the prey items, inferred from skeletal elements, suggests that the bones were accumulated by a species of fish owl (*Bubo* sp. syn. *Ketupa* sp.) ([Zukswert 2014](#)).

Prehistoric pellet deposition from generations of owls using this particular roost could also have been a pathway for the introduction of bones into the archaeological deposits of the site. However, inspection of the distribution of the small vertebrate component of assemblages recovered in the three excavated trenches from Hang Trống indicated that the majority was concentrated in the uppermost excavated contexts (table 5). We conclude that the small mammal component of the upper archaeological contexts is most likely the result of mixing between modern and Late Pleistocene material and, therefore, we used their presence as a taphonomic marker and proxy for disturbance. Notably, such evidence was mostly confined to the upper few centimetres of deposit indicative both, that contexts below these carried no more than minor mixing through this particular non-anthropogenic vector, and that the roosting site itself may be of comparatively recent age – though still conceivably hundreds if not thousands of years in duration.

##### 4.4.3 Molluscs

A stand-out feature of many Mainland Southeast Asian archaeological assemblages dating from the Late- to early Post-Pleistocene are riverine or terrestrial shell middens ([Rabett et al. 2011](#)). To explore the taxonomic diversity and frequencies of represented molluscs within the Hang Trống midden, with respect to changes in site formation and wider environmental

conditions, a column was removed (20 x 10 cm) from the north-facing section of Trench 1 to a depth of 1.33 m below the current ground surface (figure 5).

The majority of the species identified were terrestrial molluscs (table 6). This included three species from the genus *Cyclophorus*: *Cyclophorus theodori*, *Cyclophorus unicus* and *Cyclophorus* cf. *cambodjiensis* – all of which live on or around trees. In addition to *Cyclophorus*, other terrestrial species were also identified, namely *Cryptozona* cf. *chrysoraphe*, and Zonitidae cf. *Oxychilus* sp. A further two examples of gastropods were noted for which only genus level attributions could be made; *Camaena* sp. and *Amphidromus* sp. both favour trees and shrub habitats.

Little data is available in the literature for the habitats of *Cryptozona* cf. *chrysoraphe*, and Zonitidae cf. *Oxychilus* sp. other than that the genera concerned have affinities with woodland settings. They continue to be found locally. A small number of aquatic taxa were also identified: *Unio* spp. and *Planorbis* sp. both favour larger bodies of freshwater such as lakes, ponds and rivers; *Cerithrium* sp. and *Ellobium aurisjudae* can occupy a broad range of aquatic/semi-aquatic habitats including estuaries, mangroves, shallow water and rivers. Minimum Number of Individuals (MNI) counts were calculated on the basis of the largest number of identifiable parts present per sample (i.e. whole shells, with lip or tip intact; plus the highest number of fragmented lips or tips). Total unburnt and burnt shell MNI counts, plotted by depth, are presented for comparison against vertebrate faunal remains from Trench 1 in figure 5.

#### << Table 6 mollusc >>

#### 4.5 Technological evidence

Typological classification of Southeast Asian lithic assemblages has been hampered by the often amorphous character of the material. While types can and have been identified, the utility of this approach to Hoabinhian or Hoabinhian-like technologies has come into question (e.g. Reynolds 1990; Marwick 2008, Forestier *et al.* 2013). In an attempt to help address this issue, and to foster greater integration between technological actions and resource acquisition, the Hang Trống lithic assemblage was studied using techno-typological (see Reynolds 1989) and attribute analysis (see Tostevin 2013) methodologies. Both concentrate on the way that lithics were manufactured and explore variability within assemblages rather than categorising them according to strict types. The former method emphasises distinguishing broad technological classes that group generally similar pieces; the latter emphasises the dynamic nature of the lithic reduction process and how that variability can be understood in relation to human adaptive strategies. The use of two complementary methods, applied by two different researchers, also helped us to control for biases in distinguishing artefacts from non-artefactual pieces. In what is an overwhelmingly coarse-grained assemblage, often without prominent bulbs of percussion or other clear markers of technological practice, we accepted that levels of observer variance may exceed 5-10% (after Proffitt & de la Torre 2014).

An initial techno-typological study of selected pieces from the 2009 excavations at the site (context 8016, Trench 1) has been published elsewhere (see Rabett 2012, figure 6:17). This material has since been returned to Vietnam and is not integrated into the current study as it was not feasible to conduct an attribute-based analysis of it. Briefly, this initial study identified a core rejuvenation flake with at least two or three previous hard-hammer flake removals; a blade-flake, possibly struck using a soft-hammer technique; a large, thick

secondary flake with a plain platform – this piece had slight rounding along the distal edge and evidence on its dorsal surface that many small flakes had been removed prior to this one – and finally a ‘steep-edged’ core that may also have been used as a pounder. These pieces suggested on-site production, including basic core preparation and possible use of soft-hammer reduction. Both study samples covered herein are drawn from the most extensive period of excavations at the site (in 2010). Material studied consists of, that recovered from LGM deposits in the lower half of Trench 1 (8201-8111) and post-LGM (MIS-2) deposits (and likely post-LGM deposits) from trenches 3 and 2, respectively.

Techno-typological analysis identified  $n = 74$  modified pieces (table 7), with a further 13 of ambiguous classification. The industry to be based on single-platformed hard hammer reduction, with no evidence for core preparation, platform edge maintenance or platform preparation. There are a small number of silet flakes that may derive from the use of larger or heavier hammers (silet flakes are split longitudinally through the bulb of percussion). Secondary and tertiary flakes dominate the sample, indicative of non-intensive use of materials and possibly the reduction of materials elsewhere (either on site or off site). The formal flake tool-type list is limited to two scrapers and a burin (table 7, artefacts 11 [from cleaning], 17 & 42, respectively). The pebble-based portion of the inventory includes a hammer/pounder with patches of hammering, pitting and micro-flaking on both ends. Seven pieces were defined as ‘short-axe’ fragments (*see* figure 6), following the northern Vietnamese typological nomenclature for the Hoabinhian. Attribute analysis of the sample (Phan 2014, Appendix B) identified  $n = 76$  modified pieces (with 63 confidently defined as cores, pebble tools, flakes, retouched flakes or flake fragments) (table 7). The sample may be regarded as the background debris of lithic tool assisted activities. In-situ knapping could not, though, be further demonstrated and refits were not identified.

<< Table 7 lithics >>

<< Figure 6 illustration of ‘short-axes’ from Hang Trống >>

## 5. DISCUSSION

Distinct patterning within the results of the various proxy analyses permitted separation into four depositional phases: a mixed phase and three largely undisturbed phases of site use. Phases are described from the base of excavation upwards.

### 5.1 Phase III

This phase incorporates material from contexts identified between the current excavation base, dated to 24,244-24,887 cal. BP (UBA-17272), and the appearance of a large shell midden (commencing in Phase II). Data are drawn exclusively from Trench 1 (contexts 8100-8111) (figure 5). Compared to contexts higher in the stratigraphic sequence, sediments at this depth are almost devoid of cyclophorids; however, infrequent freshwater shells, charcoal and small numbers of bone fragments and cultural material was recovered to full depth. Notable amongst the modest vertebrate fauna recovered (table 5) was a complete metapodial from a *Manis* sp. (pangolin) in context (8111). This piece is slightly damaged on the dorsal distal surface (lateral side) and proximally on the lateral side, preventing clear identification to element. Two species of pangolin are presently found in Vietnam: *Manis pentadactyla* (Chinese pangolin) in the north and centre of the country, and *M. javanica* (Sunda pangolin) in the centre and south (Lekagul & McNeely 1988; Sterling *et al.* 2006).

Techno-typological analysis produced a diverse range of pieces, a greater predominance of which was made on limestone (62.5%,  $n = 10/16$  – single surface find excluded), including: one core, five flakes, one flake fragment (from section cleaning), one pebble fragment/pebble tool ('short axe') and a burnt limestone pebble fragment. Shale pieces were less common (25%  $n = 4/16$ ): two flakes and two pebble fragments ('short-axes'). This phase also yielded the site's only quartzite piece (artefact 17) – a tertiary, retouched flake ('scraper'), and the only piece of sandstone: a flaked pebble fragment. Fourteen modified pieces were also identified during the attribute analysis. These comprised: 1 core, 3 pebble tools, 7 complete flakes, 1 flake fragment 1 one retouched flake.

As this phase broadly covers the LGM it is significant that fragments of Dipterocarpaceae (cf. *Dipterocarpus*) and Sapotaceae (cf. *Manilkara* or *Palaquium*) were found in contexts (8100, 8103 & 8110, & in 8105) (table 2). The genus *Dipterocarpus* contains c. 70 species of large (c. 60 m) rainforest trees, with a wide geographic distribution through the tropics (see Ashton 1988). Sapotaceae (cf. *Manilkara* sp. or *Palaquium* sp.) is a pantropical family of flowering plants, most species of which produce edible fruit. Pinaceae and two endocarps of *Celtis* sp. (Family: Ulmaceae) were also recovered. The former ranges from subarctic to tropical environments, including limestone karst (Sterling *et al.* 2006) and also yields edible seeds. On the basis of the available comparatives, the Hang Trông *Celtis* sp. specimens are more likely to be *C. australis* L. (Mediterranean hackberry) or less likely the *C. elim. sinensis* Piers (Chinese hackberry). *C. australis* is a component of more open dry deciduous forest regionally (e.g. Sun *et al.* 1986).

The  $n$ -alkane data from this depth in the excavation indicates a vegetation cover that is similar to today (figure 4). The plants around the cave do not seem to have been adversely effected by the climate downturn. Compound-specific isotope analysis of both  $n$ -alkanoic acid and  $n$ -alkane higher plant lipids also indicate that the prevailing occurrence of  $C_3$  vegetation. Pollen records from the South China Sea during this time indicate herbs and grasses, with some temperate broad leaf plant types dominated most lowland landscapes (Sun *et al.* 2003). The well-watered and enclosed environment within the Trảng An massif appears to have provided a stable set of conditions for sub-tropical plants despite the climatic downturn.

## 5.2 Phase II

The deposits Phase II span a period of approximately 2000-2500 years: from 17,845-18,520 (UBA-14885) and 18,229-18,708 cal. BP (UBA-14886) to a little before 15,522-16,663 cal. BP (UBA-14884). Macro-vertebrate faunal remains are comparatively rare (NISP  $n = 57$ ), with indeterminate large mammal fragments making up the bulk of this total (table 5). Deer (Cervidae) are represented only by distal elements (including a burnt phalange and carpal). Cercopithecidae also feature, with notable finds including two mandibular fragments with partial *in-situ* dentition. The first of these was recovered from near the base of the shell midden in Trench 1 context (8016). Charcoal from this context is dated to 18,229-18,708 cal. BP. The specimen consists of a right  $M_1$  with alveoli for  $M_2$  and  $M_3$  – indicating an adult of five or more years of age (Smith *et al.* 1994) and is attributable to *Macaca* cf. *mulatta* based on biometric data (Swindler 2002, table 112). The second, recovered from Trench 3 context (8312), was from a juvenile: comprising a (left)  $dp_1$  in wear a  $dp_2$  and an erupting  $M_1$  in one fragment, with a similarly faceted  $dp_2$  in a second associated fragment (from the right of the mandible). Evidence of other terrestrial-arboreal and arboreal taxa includes: Sciuridae, from mandible fragments equivalent in size to *Sciurus* sp., Hystricidae by isolated teeth, and one

avian family, Phasianidae, from a fragment of left coracoid, identified as being from a gallopheasant (*Lophura* sp.) also recovered from near the base of the shell midden in Trench 1 context (8107). Micro-vertebrate remains are noticeably rare compared to higher in the sequence, where they are linked to cave-roosting birds. Infrequent freshwater crab chelae were also identified.

The lower part of this phase yielded molluscan evidence from four consecutive and approximately equal spits within context (8019) and the highest quantities (MNI) of molluscs in the phase (table 6). Cyclophorids dominated but a number of other taxa were also represented. These too are associated with forested/arboreal environments, such as *Amphidromus* sp., which lives on trees and shrubs. Three fragments of the freshwater mussel *Unio* spp. appear at this point, as do fragments of *Ellobium aurisjudae*. Unionidae generally inhabit freshwater environments, such as slow rivers, lakes or ponds; while *Ellobium aurisjudae* can inhabit quite a broad range of aquatic environments, from muddy estuaries to mangroves (Nguyen Ngoc Thach 2005). From the top of the phase (contexts 8010 & 8011) only a low quantity of molluscs were recovered and three species of *Cyclophorus*. The molluscan data from Phase II, though, supports the picture that the area around the cave was at this time forested and contained a source of freshwater in its vicinity.

Techno-typological analysis indicated that shale was the predominant raw material used in the manufacturing of lithic artefacts 89.6 per cent ( $n = 26/29$  – excluding potentially natural pieces). This included flakes, flake fragments and one pebble fragment ('short axe'). The three remaining pieces were limestone flakes/flake fragments (table 7). Attribute analysis identifications were more conservative:  $n = 20$ , comprising: 1 pebble tool, 9 complete flakes and 10 flake fragments.

Although palynological work was hampered by the degraded state of pollen grains, some identification was nonetheless possible for Phase II, including tree families: Brassicaceae (Crucifer), Rubiaceae, Leguminosae (Fabaceae/Mimosa), Betulaceae and *Pinus* or *Podocarpus*; and more diverse families, such as Moraceae or Urticaceae, which are both mixed families of trees, herbs, climbers or succulents, preferring warm tropical and lowland forest regions; Ulmaceae trees and shrubs persist, and Poaceae, which is an open ground grass taxa also appears. Proximal pollen records from comparable conditions within China reflect the findings from Hang Trống, whereby sub-tropical forests dominate (Wang *et al.* 2012). Charred fragments of Dipterocarpaceae (cf. *Dipterocarpus*) were recovered from contexts (8016 & 8100); as were specimens of Pinaceae and Sapotaceae (context 8100). Fragments of Rhizophoraceae (cf. *Rhizophora*) tree and shrub taxa came from the top of the phase (context 8010) in Trench 1. These latter are significant components of mangrove forest (Nguyen Ngoc Chinh *et al.* 1996) though the potential proximity of the site to this habitat at this time has yet to be determined. Leguminosae (cf. *Dialium* sp.) fragments were identified from this phase in Trench 3, context (8308). Remains of *Celtis* sp. endocarps are notably increased from Phase III, while those of the nut *Canarium* spp. (Family: Burseraceae) appear for the first time (table 2). It is unclear if these changes relate to more intensive use of the site during Phase II (see figure 5) or to a shift in (through continuing presence of) forest conditions.

Towards the base of this phase (141.0 m. asl.) the *n*-alkane data display a slight increase in average chain length. This potentially signals that a resurgent southwest monsoon during the early post-LGM provided a more favourable growing environment; something that may also have been linked to more intensive use of the site. Compound-specific isotope analysis of both *n*-alkanoic acid and *n*-alkane higher plant lipids remains



stable through-out the phase, indicating a predominantly C<sub>3</sub> vegetation cover, contradictory to evidence from islands further south where C<sub>4</sub> plant-type dominate at this time (Wurster *et al.* 2010). The *n*-alkane data indicate vegetation broadly similar to that seen in the property today.

### 5.3 Phase I

The final intact occupational phase in the excavation is associated with an age 14,056-14,902 cal. BP (UBA-21288) from Trench 3 and 15,522-16,663 cal. BP (UBA-14884) in Trench 1. Representatives from five macro-vertebrate mammalian families were identified: Cervidae, Hystricidae, Mustelidae (*Arctonyx collaris*), Cercopithidae (cf. *Macaca* sp.) and Sciuridae. The deer (Cervidae) remains were represented by elements from the distal appendicular skeleton, namely phalanges and a left tarsal. Evidence of porcupines (Hystricidae) comprised of isolated teeth; the vector of introduction for this animal is unclear as porcupines are known to frequent caves in this region (e.g. Bacon *et al.* 2008). A well preserved right distal humerus belonging to a hog badger (*Arctonyx collaris*), provides one of the few intact element recoveries from Trench 1. A discrete assemblage of bone fragments from this phase in Trench 3 included pieces of humerus, ulna and femur from a monkey (cf. *Macaca* sp.), together with a complete femur and burnt mandible from a squirrel (Sciuridae). This concentration of remains – unique on the site – was found within context (8308) (sq. 610/715) and dated 14,056-14,902 cal. BP (UBA-21288). The faunal remains were also found in association with three stone implements, with patches of charcoal and possibly ochre adhering to their surfaces (table 7, artefacts 48 & 52) and several flakes, all clustered underneath a single whole large freshwater mussel. This context was excavated beneath a large slab of roof fall, which also carried red and black linear pigmentation on its western face: it is as yet unclear if this colouration was deliberately applied or was the result of burning from hearths in the immediate vicinity, though the concentrated nature of these finds certainly bears the hallmarks of having been deliberately placed.

This phase also included low frequencies of micro-vertebrate remains. These were dark in colour (compared to examples from the taphonomic control and mixed contexts above) with calcium carbonate deposition clearly evident. Included among these remains were bones of birds, fish, rodents and shrews. The presence of this material is likely to be an indication that owls were roosting in this part of the cave; though the filtering of micro-vertebrate bones through the coarse matrix of shells cannot be completely discounted. The bones of insectivorous bats are also present, including *Taphozous* sp. (tomb bats). These are also interpreted as cave-dwelling fauna and, therefore, almost certainly natural introductions. Small quantities of crab chelae were again found, including burnt specimens. These are more reliably considered to have been introduced to the site by people.

Data from the molluscan column sample from Trench 1 (table 6) that relates to the lower part of this phase includes material from three contexts (8007a-c & 8009), which is again dominated by *Cyclophorus* sp., *Cryptozona* cf. *chrysoraphe* and Zonitidae cf. *Oxychilus* sp., present also now is *Camaena* sp. – an herbivorous terrestrial mollusc that lives in and around shrubs and trees. This taxonomic representation is again indicative of a forested local environment. Single pieces of the aquatic *Unio* spp. and *Cerithrium* sp. were also identified; indications that nearby freshwater habitats were being exploited. Burnt shell suggests hearths were set in the middle of the cave. The molluscan sample from the upper part of this phase (contexts 8006, 8005, 8004, 8003b & 8003a) was found to be dominated by *Cyclophorus* sp. Quantities of *Cryptozona* cf. *chrysoraphe* and Zonitidae cf. *Oxychilus* sp.

were also identified. *Camaena* sp., which is found deeper in this part of the sequence were not present. This may indicate a shift in local conditions or a change in human gathering patterns.

Techno-typological analysis of the lithic assemblage from the 2010 excavations identified  $n = 21$  artefacts from Phase I in Trenches 2 and 3 (table 7). This consisted a predominance of shale flakes, flake fragments and blade/flake (with one a burin) specimens; as well as a shale pebble chopper/core ('chopper') and retouched pebble fragment ('short-axe') – 71.4% per cent ( $n = 15/21$  – excluding potentially natural pieces) of the total. Other modified pieces were flakes, core and pebble ('pounder'), all of limestone; and a burnt pebble fragment of unknown material. A lower total of artefacts ( $n = 17$  modified pieces) was identified through the attribute analysis study, comprising 10 complete flakes, 3 flake fragments, three pebble tools and a core were identified.

The climate is likely to have been variable during this phase: insolation levels were increasing; however the phase is also punctuated by H1, a noticeably cold-dry period, indicated by sea surface records from the Gulf of Tonkin and the Hulu Cave speleothem record (e.g. Huang 1997; Pelejero *et al.* 1999; Wang *et al.* 2001; Visser *et al.* 2003; Oppo & Sun 2005; Wei *et al.* 2007; Shintani *et al.* 2011). Pollen recovered from within this phase was noticeably degraded, though appears synchronous with a proximal lake pollen record from southern China (Wang *et al.* 2012). Grains of Brassicaceae (Crucifer) and Rubiaceae were, also identified, as were pollen grains of Hamamelidaceae, a family formed of tropical trees and shrubs. Betulaceae and Cyperaceae represent other community components from this phase. Dipterocarpaceae were confirmed in the macro-botanical data (table 2). At this time the coastline is thought to have been still up to 500 km from the Trảng An massif (Yao *et al.* 2009). Sedimentary facies from coring in the Song Hong Delta (core ND-1, 20.372778 N, 106.146667 E) show that *c.* 14,900 cal. BP the low-lying landscape beyond the massif contained channels from a meandering river system (Tanabe *et al.* 2006).

Close to the onset of this phase (141.7 m asl) *n*-alkane chain lengths are noticeably shorter (average *n*-alkane chain length C<sub>29</sub>). It is hypothesized that this reflects the change to colder and dryer conditions in H1. Taken together with increased post-glacial insolation, these factors inhibited production of the protective long-chain *n*-alkanes needed in this sub-tropical environment; plants only had sufficient resources to maintain shorter *n*-alkanes. Towards the end of the phase, and the expected increase in temperature and moisture, longer *n*-alkane lengths were being produced by the vegetation. Compound-specific isotope analysis of *n*-alkanoic acid and *n*-alkane higher plant lipids from Hang Trống sediments all indicate that the interior of the massif was still characterized by a dominant C<sub>3</sub> vegetation signal that was distinct from the environment setting that lay beyond it.

#### 5.4 Mixed Phase

Surface contexts (corresponding to approximately the upper-most 10 cm in each trench) were allocated to a mixed phase due to the likelihood of modern contamination of the sediments. This is borne out through the two radiocarbon dates obtained: 12,594-12,854 cal. BP (UBA-09301) from our reconnaissance test-pitting here in 2007 and a contrasting date of 32-355 cal. BP (UBA-14883) from context (8001) in Trench 1 in 2009. In faunal terms, these contexts are dominated (NISP) by the remains of rodents (Muridae), amphibians (Ranidae: *Rana* sp.), bird bones (small passerines) and small numbers of fish bones. All are considered most likely to derive from owl pellets from roosting sites in the roof of the cave. Bones of insectivorous bats (Chiroptera) and reptiles (*cf.* Gekkonidae) likely represent natural deaths

of cave-dwelling fauna. The bones of larger mammals were rare in this phase, but included a fragment of vertebrae from a non-human primate, possibly relating to historic occasional use of the site by hunters.

Molluscan evidence is relatively diverse, including *Clausilia* cf. *proctostoma* – a species that was only otherwise observed during the LGM (*see* Phase III), which may point to a cooling from within the Pleistocene data of these mixed deposits. Aquatic taxa are present, with *Planorbis* sp. a species of freshwater snail and a spike in the percentage quantities of *Cyclophorus*, from c. 65 per cent of the sample to 85 per cent in little more than a couple of centimetres. All of the terrestrial species identified from this phase are known to live in, around or beneath trees, and in the case of other identifications, such as *Cryptozonia* cf. *chrysoraphe* and Zonitidae cf. *Oxychilus* sp., they are likely to inhabit very broad environmental spheres that encompass arboreal environments.

Technologically, the upper-most contexts (8200-8203) in Trench 2 and (8300-8302) in Trench 3 produced eight lithics through techno-typological analysis, and eight through attribute analysis. Techno-typologically, the identified pieces included 1 limestone pebble fragment that might have been a core, 3 shale flake fragments, 3 limestone flakes and a limestone pebble flaked fragment ('short axe') (table 7). Attribute analysis of the same sample identified 1 core, 5 complete flakes, 1 pebble tool ('short-axe') and 1 flake fragment.

This phase contains the only substantial concentration (60%) of identifiable pollen recovered. This is almost certainly due to modern wind-blown grains becoming trapped within the sediments as erosion and localized micro-re-deposition of these takes place, given the cave's morphology. Charred fragments from samples taken within this phase reveal a number of Dipterocarpaceae specimens; together with *Celtis* sp. and *Canarium* spp. fragments (table 2). The climate signal contains both evidence of warm, moist conditions and also indications of sediment accumulating under cooler conditions. *n*-Alkane biomarkers average high chain lengths, possibly simply a reflection of contemporary sub-tropical conditions.

## 6. CONCLUSION

The results of work conducted at Hang Trống highlight the potential resilience of limestone forest habitats in the northern tropical zone of mainland Southeast Asia. Multiple lines of evidence suggest that these forest formations can sustain themselves through substantial local climatic and environmental change.

Palynological and macro-botanic evidence indicates the persistence of tropical forest cover within Trảng An during the LGM, a period when the surrounding lowlands were characterised by grassland environments. Invertebrate and vertebrate fauna data corroborate the strong arboreal nature of the local habitats within the massif and their durability across this climate downturn. Compound specific analysis of plant biomarkers shows a common C<sub>3</sub> photosynthetic pathway for local vegetation. The *n*-alkanes recovered from Hang Trống indicate that the most dominant chain length in the modern vegetation (C<sub>31</sub>) is also predominant in the sedimentary record at the site. This is concurrent with other plant zones from this latitude, thus we can be confident that evidence from our analysis of *n*-alkanes reflects, with some clarity, regional climatic conditions. At c. 15 kBP a change in the forest dynamics is indicated by a reduction in C<sub>31</sub> *n*-alkane and an increase in C<sub>27</sub> *n*-alkane. A dramatic change in *n*-alkane chain length at c. 15 kBP coincides with the H1 event. This can be attributed to a strengthening of the northeast monsoon during this interval (e.g. [Huang et](#)

*al.* 2011). The associated increase in aridity does not, however, appear to have instigated a substantial shift in local vegetation; conceivably in part due to adaptations toward aridity within limestone forest, conditions rebounded and have persisted through the Holocene.

The occupational record from Hang Trống appears to indicate intervals of more and less intensive site-use, peaking in the period immediately after the LGM, but always at a relatively low scale compared, for example to the later midden deposits from elsewhere in the massif, such as Hang Boi (Rabett *et al.* 2011). The abundance of particularly cyclophorids during the wet southwest monsoon is taken as a likely indicator that this was probably the main season of use, though visits outside of this time cannot be ruled out. There are hints of caching activity that, if correct, are suggestive of provisioning. The nature of this through-cave certainly raises the possibility that it was utilised primarily as a corridor to travel between valleys; the generally equitable conditions providing a useful temporary stopping point.

The technological sample obtained from the 2010 excavation was small ( $n = 74$  artefacts), precluding statistically robust statements about observed variability, hence the conclusions we present here are tentative. Patterning in the frequency of artefact categories, form of platform (i.e. cortical, plain, pointed, crushed), flaking direction, striking platform thickness and exterior platform angle) appeared within the lithic assemblage suggests a shift in knapping behaviour and hence potentially mobility between the LGM and the post-LGM (Phan 2014). This shift was from one where limestone predominated, but where a wider range of materials were utilised, during the Phase III (LGM) use of the site, to sourcing that more particularly focused on shale, especially in Phase II. A higher observed frequency of dorsal flake scar counts is also suggestive of greater reduction intensity during the LGM compared to later phases of site-use. The frequency of flakes and flake fragments is considerably lower than is seen during the later phases. Both raise the possibility that occupation of the site during this period involved a change in curation and ‘technological organization’ (after Nelson 1991), if not in site-use, and may turn out to reflect difference in risk levels and mobility strategies (Phan 2014). As the classes of artefact recovered appear to have been largely consistent through-out, variability was perhaps situated primarily in how technological strategies were implemented rather than in technological forms themselves. In other words: the general reduction strategy (the technology) appears to not have changed, but attribute analysis of our sample suggests that components or steps in that strategy (the technological organization) were altered in response to external pressures.

Pieces resembling Hoabinhian ‘short-axes’ persisted through all three phases (table 7, artefacts 1, 7, 10, 13, 29, 46 & 87). Artefact 13 from this list is from a secure context overlying the deepest current date from Hang Trống (*c.* 24.4 cal. kBP), indicating an age-range for this artefact-type from close to that of the earliest proposed for the Hoabinhian in Vietnam (*cf.* Yi *et al.* 2008) up to the top of the sequence. This finding is in-keeping with the technological durability that has been observed elsewhere for the Hoabinhian (e.g. Forestier *et al.* 2015). The lack of other supposed type-artefacts attributed to the Hoabinhian, most notably ‘Sumatraliths’ or unifacial discoid pieces, which have been considered definitive elements (e.g. Ha Van Tan 1997; White & Gorman 2004) also highlights the variable expressions that the Hoabinhian encompasses. In a reversion to the original adaptive emphasis of Gorman (1970, 1971), we consider that the Hoabinhian record might be more aptly described in terms of a technological (and subsistence) ‘eco-system’ of responses rather than as a singular entity (see also e.g. Nguyen Viet 2000; Nguyen Gia Doi 2005).

Indications from the current study suggest that foraging groups focused attention on the exploitation of resources available within the Tràng An massif at least on a seasonal basis over a long period of time: many thousands of years. This stability and concentration of resources carries wider implications for our understanding of the archaeology of the region. The consistent recovery of occupational evidence from limestone environments may have significance beyond the better preservation offered by these settings and traditional research focus within them. They may represent genuine ‘islands of humanity’: points on the landscape that early populations occupied consistently and exploited preferentially over millennia because of the surety they offered in resource access (Rabett *et al.* 2011), a fact likely tied to their refugial character (see Phoca-Cosmetatou & Rabett 2014).

The results of our analysis point strongly to the presence of a forested environment within Tràng An throughout the course of the last twelve thousand years of MIS-2, and provide testament to its resilience in the face of large-scale climate and environmental change. Forest continuity here into the Holocene and present day (see Rabett *et al.* 2011; Ludgate *et al.* forthcoming) underscores and extends that resilience. The effects of modern climate change on tropical environments had been seen as secondary to that predicted from deforestation and extractive industry (Sala *et al.* 2000); however, the projected climate-driven impact on environments and communities in Southeast Asia has grown significantly in the last decade. Coastal and low-lying island habitats and their associated populations are now known to face an acute threat from climate-driven inundation and environmental transformation (Nicholls & Cazenave 2010; Rietbroek *et al.* 2016). While the corpus of literature on the impact to coastal habitats globally gathers apace (e.g. Virah-Sawmy *et al.* 2009; Kirwan *et al.* 2010; Holland 2012; Mendoza-González *et al.* 2013; Hunter *et al.* 2015), with few recent exceptions (e.g. Wetzel *et al.* 2012; Faridah-Hanum *et al.* 2014; Latinne *et al.* 2015) the effect of climate change on coastal tropical terrestrial habitats is still an under-represented field, and a body of research to which this study contributes. In closing, then we suggest that there are notable implications for conservation arising from the Hang Trống study. Firstly, long-term early human presence in Tràng An, even under conditions when vegetation was more stressed, does not appear to have been adverse to forest survival. That being so, there is an argument to be made that with careful management the impact of modern visitors to this and similar karstic properties on the World Heritage List could be sustainable. Secondly, it is now generally accepted that refugia may offer the best chance for the survival of many taxa under conditions of anthropogenic climate change (e.g. Médail & Diadema 2009; Ashcroft 2010; Keppel & Wardell-Johnson 2012; Keppel *et al.* 2012). The data from Hang Trống, therefore, add additional weight to that contention. In line with other refugial habitats, tropical limestone forests should be considered as a conservation priority. Finally, faced with the prospect of substantial transformation through climate change and regional sea-level rise our findings also give reason for optimism. With diligent management the chances of survival remain positive for these biodiverse, richly endemic and archaeologically bountiful landscapes.

## ACKNOWLEDGMENTS

We would like to thank the Tràng An Management Board, the Ninh Bình Provincial People’s Committee and Xuan Truong Enterprise for continuing to recognize the importance of archaeological research within the Tràng An World Heritage property. RR gratefully acknowledges that key funding for much of this work came through the McDonald Institute for Archaeological Research, and via the Evans Fund, Department of Social Anthropology,



both at University of Cambridge. He also wishes to further thank the McDonald Institute, and the School of Geography, Archaeology and Palaeoecology (including the 14Chrono Centre) at Queen's University, Belfast for their long-term involvement with this project. The authors are very grateful to Patrick Roberts (associate editor) for the opportunity to contribute to this timely and important special issue; and for the constructive comments that he and three anonymous reviewers made on the manuscript.

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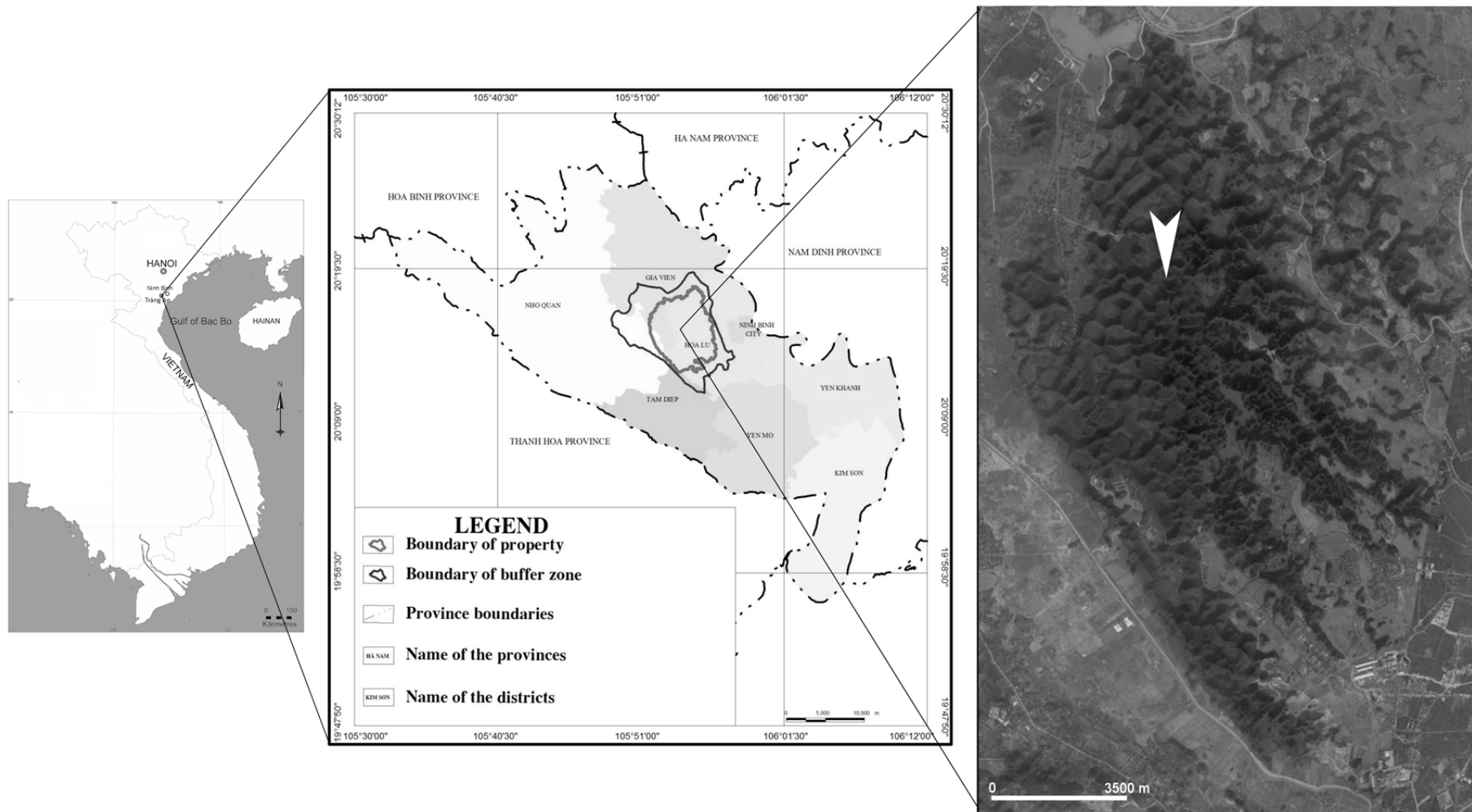


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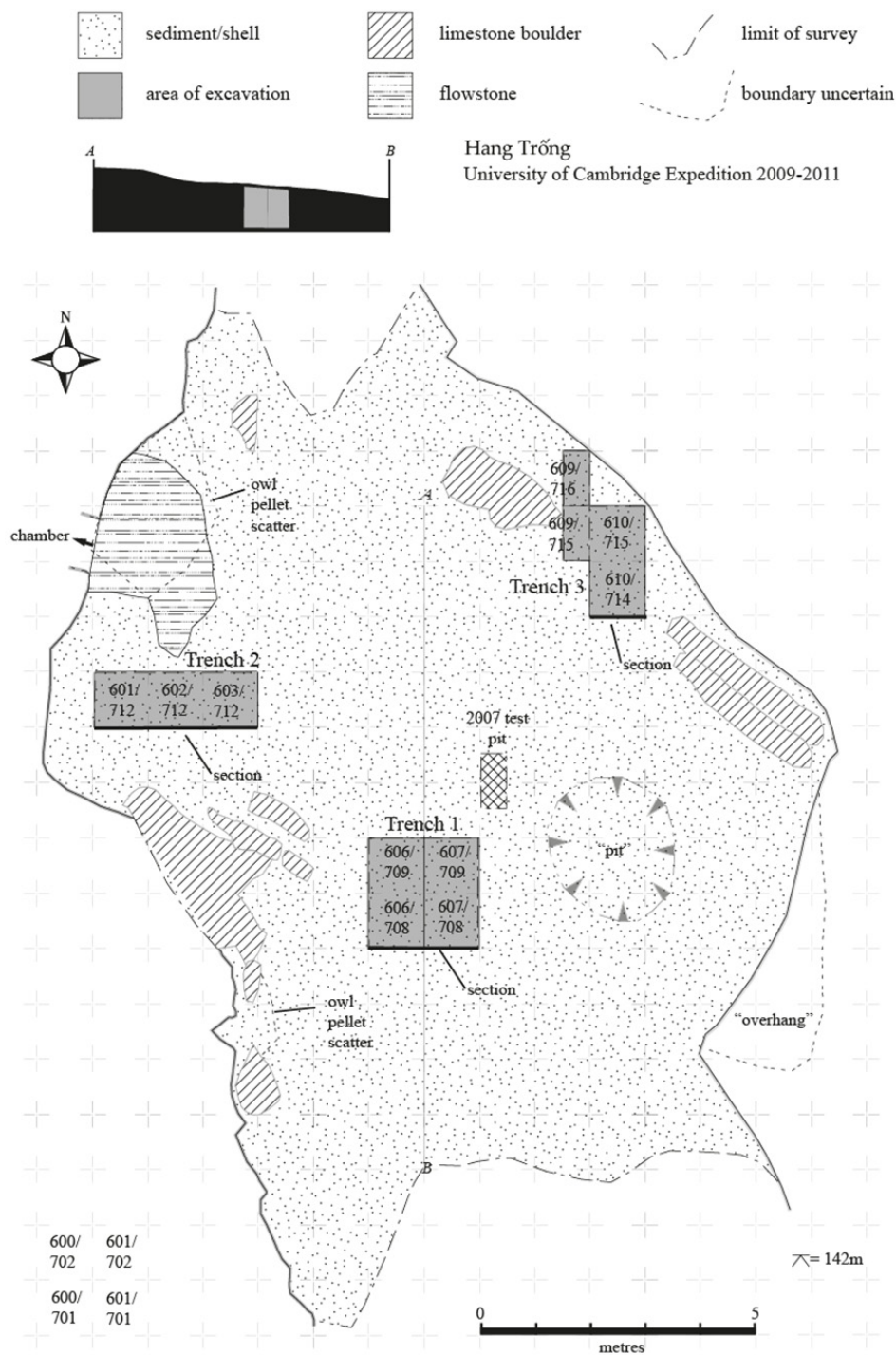
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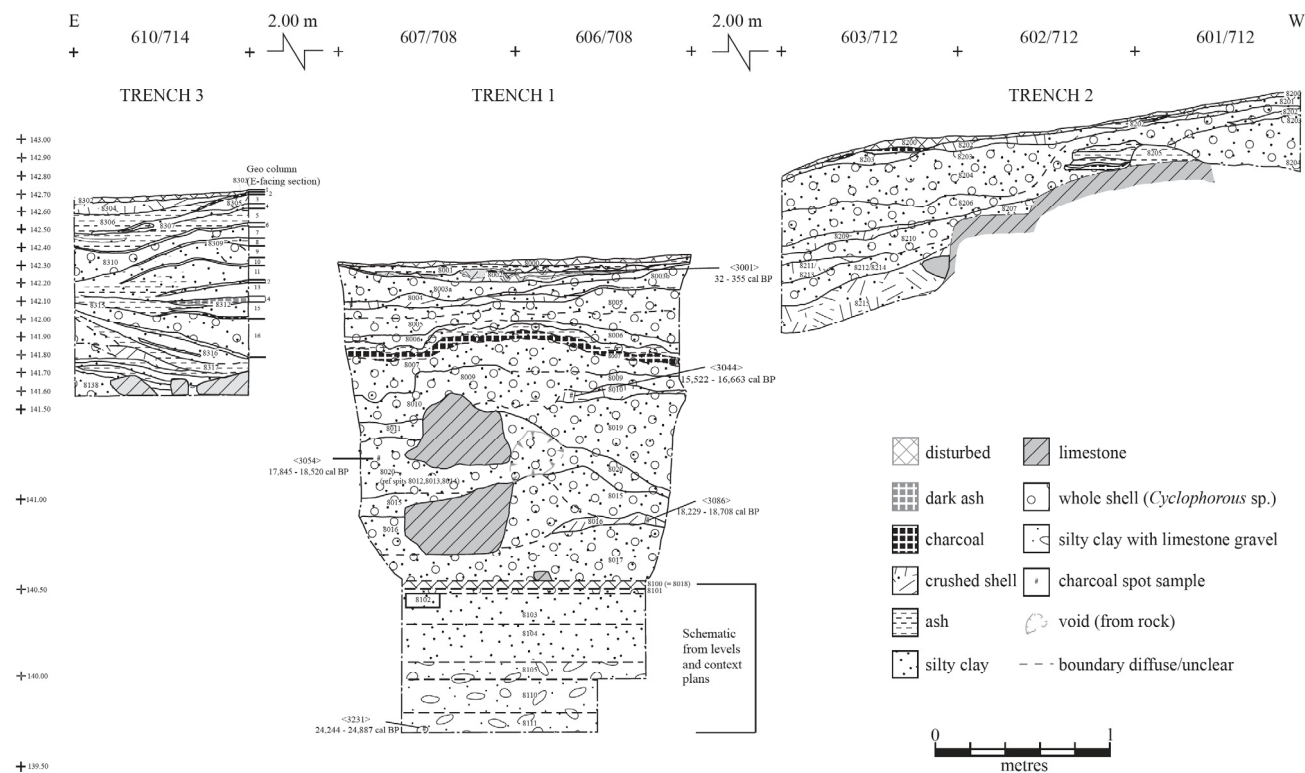


<< Figure 1, the location of Tràng An, Ninh Binh province; the World Heritage Site core and buffer zone boundaries of the property; and the location of Hang Trống within the Tràng An limestone massif (right-hand image: Google Earth). >>

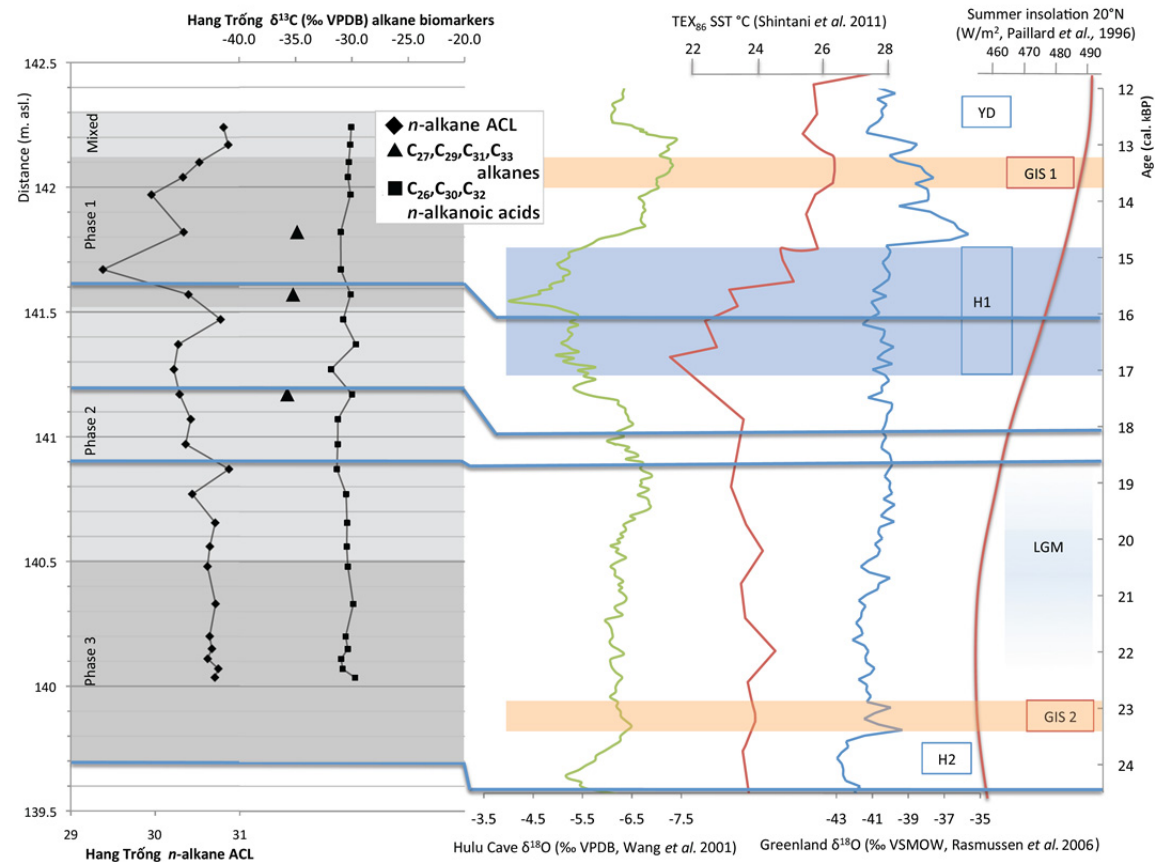




<< Figure 2, plan of Hang Trống showing pertinent site features, including geoarchaeological sampling points, the locations of the 2007 test pit and of TR1-3. Survey: J. Hawkes, C. Stimpson & M. Verhoeven. Drawing: C. Stimpson. >>



<< Figure 3, composite of the north-facing sections from Trenches 1-3 across Hang Trống. East-West distance between trenches has been fore-shortened for the purpose of this illustration. Original drawings: Borbala Nyiri, Lucy Farr, Mike Morley, Jason Hawkes, Marc Verhoeven, David Simpson and Jo Appleby; composition: C. Stimpson. >>



<< Figure 4, *n*-Alkane data from Hang Trống compared to published palaeoarchive data during the LGM: thick horizontal lines indicate <sup>14</sup>C dates from Hang Trống at respective depths and ages along the cal. kBP scale. Hang Trống  $\delta^{13}\text{C}$  biomarkers display the consistent trend of C<sub>3</sub> vegetation surrounding the site. Oxygen isotopes from Hulu cave have been shown to provide evidence of monsoonal changes, which coincide with sea surface temperature changes (SST) in the China Sea. Oxygen isotope records (Rasmussen *et al.* 2006) indicate global temperature changes, specifically Greenland Interstadial 1 (GIS 1) and GIS 2. Both oxygen isotope and SST records indicate little change during the LGM, during which time summer insolation values were low (c. 460W/m<sup>2</sup>). *n*-Alkane average chain length values (ACL) from Hang Trống also display a consistent trend across the LGM. H1 (c. 15 kBP), however, sees a dramatic change: a drop in oxygen isotope values from Hulu Cave, SST from the South China Sea and ACL from Hang Trống. Drawing: N. Ludgate >>

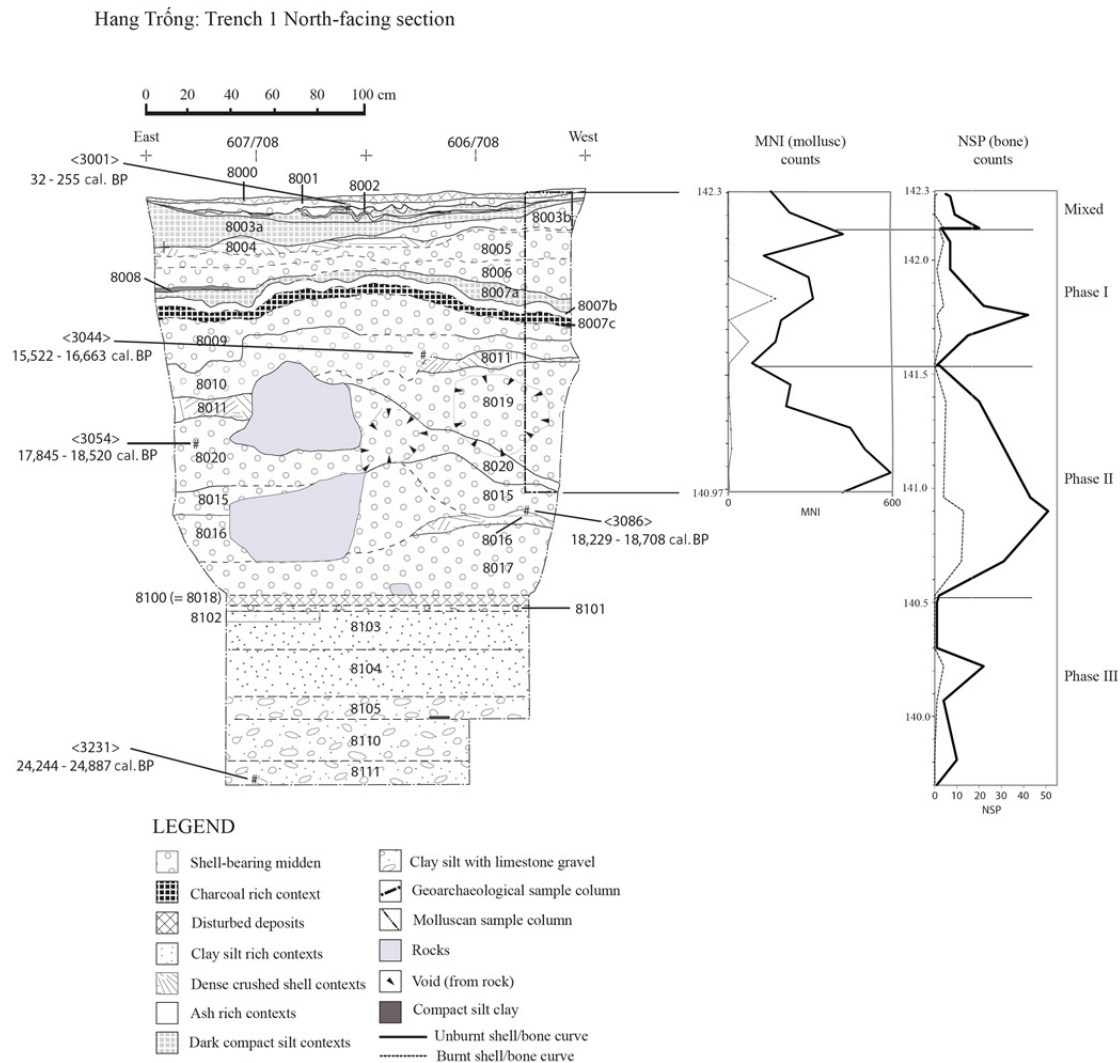
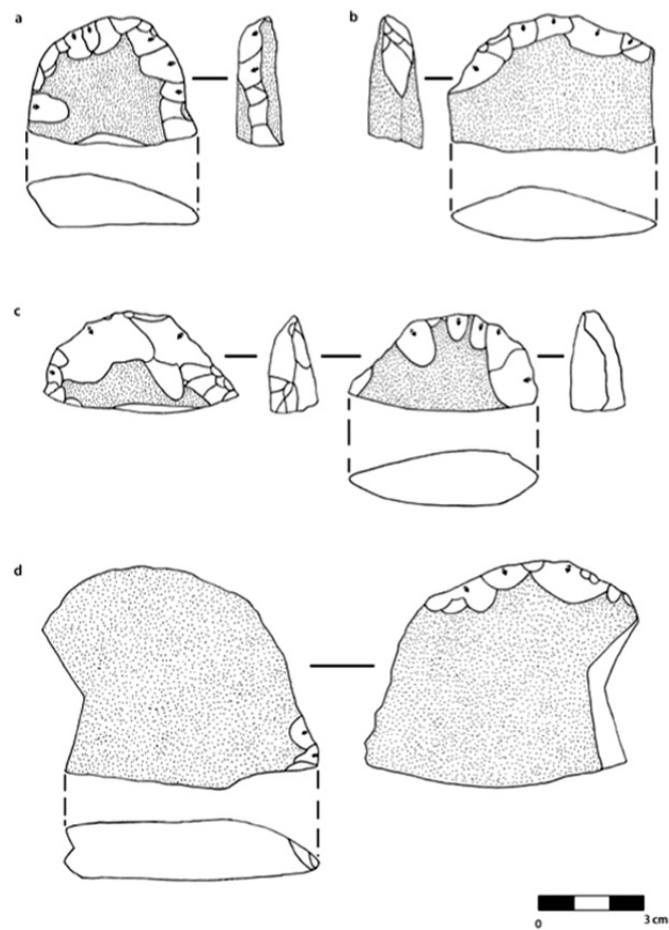


Figure 5, north-facing section of Trench 1 at Hang Trống, showing molluscan sample columns, molluscan (MNI) and vertebrate faunal (NSP) count curves (latter combining data from all three excavated trenches). Drawing: C. Stimpson and R. Rabett. >>



<< Figure 6, examples of flaked pebble tools or 'short-axes' from Hang Trống: Artefact nos. 13 (a), 7 (b), 10 (c) and 46 (d) – see table 7. Drawing: L. Phan. >>

<b>Grid Square</b>	<b>Context</b>	<b>Depth (cm) relative to datum</b>	<b>Material</b>	<b>Radiocarbon years BP</b>	<b>Calibrated years BP</b>	<b>Lab code</b>	<b>Year</b>
607/708	8001	142.25	Charcoal	80±17	32-355	UBA-14883	2009
Test pit	Layer 3	Surface -3 cm*	Charcoal	10,827±32	12,594-12,854	UBA-09301	2007
610/715	8308	142.00	Charcoal	12,363±51	14,056-14,902	UBA-21288	2010
606/708	8010	141.57	Charcoal	13,227±35	15,522-16,663	UBA-14884	2009
607/708	8013(=8020)	141.22	Charcoal	14,878±45	17,845-18,520	UBA-14885	2009
606/708	8016	140.89	Charcoal	15,300±38	18,229-18,708	UBA-14886	2009
606+607/709	8111	139.71	Charcoal	20,537±61	24,244-24,887	UBA-17272	2010

<< Table 1, <sup>14</sup>C dates obtained from excavations at Hang Trống during 2007, 2009 and 2010 field-seasons; \*the shovel test pit depth measurement is approximate and not relative to site datum. All dates were calibrated using the CALIB 6.1.0 calibration curve (2-sigma). >>

Trench	Context	Phase	Sample	Canarium sp.	Celtis sp.	Dipterocarpaceae	Leguminosae	Pinaceae	Rhizophoraceae	Sapotaceae	
1	8002	Mixed	3004	2							
	Total										2
	8003		3016		2						
	8004		3006	1							
	8006		3008		2						
	8007	I	3034		15						
	8007		-	2							
	8009		3027		10						
	8010		3031					1			
	Total										33
	8012(=8020)		3052		5						
	8013(=8020)		-		5						
	8014		-		3						
	8015	II	3088		3						
	8016		-		18						
	8017		-	11							
	8100(=8018)					•		•		•	32
	Total										77
	8101		3202		1						
	8103		-		16						
	8104		3211								
	8104		-	1							
	8105	III	3213						3		
	8107		3218								
	8109		3223								
	8110		3229		2						
	Total										23
2	Mixed	8203	3415		16						
		8203	3405		1						
	Total										17
		8204	-		1						
		8206	-		1						
	I?	8207	3410					4			
		8207	-		1						
	8215	3443		2							
Total										9	
3	Mixed	8302	3600	1							
	Total										1
	I	8307	3610	1							
		8308	3619		•	•			•	100	
		8308	-	3							
	Total										104
		8309	-		2						
		8310	3616		75						
		8312	3626		13						
		8312	-	11							
		8313	-	1							
		8315	3629						19		
		8315	3628	1							
		8316	-	3							
	II	8317	3632		33						
		8317	-	2							
		8317	3638	5							
		8317	3638	1							
		8318	-	2							
		8318	3640	7							
	8318	3640	2								
	8321	-	6								
	8322	-	10								
Total										193	
Grand total										459	

<< Table 2, identified macro-botanical remains ( $n = 459$ ) from dry-sieving and flotation of sediments collected during the 2009 and 2010 excavations at Hang Trống; ‘•’ proportion of individual identified taxa not recorded for these contexts. >>



Context:	Phase	Approx. age ( <sup>14</sup> C cal. kBP)	Depth (m)	Identified grains	Unidentified grains	Degraded grains	Pollen grains counted (including unidentifiable grains):	Percentage of identified grains	Lycopodium spore counts (added for reference)
8000-8002	Mixed	0	142.25	196	77	133	329 grains	60 %	231
8009	I	15.1	141.8	10	25	145	180 grains	5.6 %	2003
8010		15.6	141.6	2	4	24	30 grains	6.7 %	1000
8016		18.5	140.9	4	10	55	69 grains	5.8 %	2008
8019	II	17.1	141.4	22	9	505	536 grains	4.1 %	513
8019		17.6	141.3	5	5	175	185 grains	2.7 %	1001
<b>Totals:</b>				<b>239</b>	<b>130</b>	<b>1037</b>	<b>1329 grains</b>		

<< Table 3, pollen counts from the pilot study at Hang Trống. >>

Context	Phase	Approx. age ( <sup>14</sup> C cal. kBP)	Dist. (m) above sea level	Apiaceae (Umbelliferae)		Aquifoliaceae (ca. Ilex )	Arecaceae	Betulaceae 3 pore (Carpinus ca.)		Betulaceae 3 pore (Corylus ca.)		Betulaceae 5 pore (ca. Alnus)		Brassicaceae	Cyperaceae	Combretaceae /Melastomataceae	Ericaceae	Euphorbiaceae (ca. Excoecaria)	Fagaceae (Castanopsis)	Hamamelidaceae	Leguminosae (Fab/ Mim)	Mimosaceae	Mimosaceae / Acacia	Moraceae/ Urticaceae	Meliaceae/ Aglaia type or Chrysophyllum type	Pinus/ Podocarpaceae	Poaceae (Gramineae)	Rosaceae	Rubiaceae	Sapotaceae	Ulmaceae	Polyad
8000-8002	Mixed	0	142.25	3	2	1			1		2	5	1	1	15		2	1	1	1	1	3 8	2	2	1 0	5	1 0	1 5				
8009	I	15.1	141.8					1		1	2						1											2			3	
8010		15.6	141.6										1						1													
8016		18.5	140.9																					2					2		2	
8019	II	17.1	141.4				4		5	8											1			1		2		2				
8019		17.6	141.3							2											1				1		1					
Identification reference:				1	1, 2, 3		1	1, 2	1	1		3			1					1, 2		1, 2	1		1							

<< Table 4, pollen identification (Trench 1) from families identified using reference collections from Zhuo Zheng (*pers. comm.*), <sup>1</sup>Tseng-Chieng (1972), <sup>2</sup>Li *et al.* (2006a) and <sup>3</sup>Institute of Botany (1982). >>

Phase	Trench	Context	Malacostraca (crab)	Osteichthyes (fish)	Amphibia (amphibians)	Reptilia (reptiles)	Aves (birds)	Mammalia (small)	Mammalia (large)	Total NISP	NSP (large animal)	Total NSP	Burnt NSP (large animal)
Mixed	1	8000	0	0	11	1	4	26	0	42	5	47	1
	1	8001	0	0	0	0	0	0	1	1	6	7	0
	1	8002	0	0	1	0	0	0	1	2	8	10	0
	2	8200	0	0	47	0	3	34	0	84	4	88	0
	2	8200-8201	0	0	1	0	1	22	2	26	4	30	1
	2	8201	0	0	1	0	5	9	1	16	11	27	0
	2	8202	0	0	0	1	1	6	0	8	1	9	0
	2	8203	2	2	1	0	0	1	1	7	35	42	1
	3	8302	0	0	0	0	0	3	0	3	1	4	0
<b>Totals</b>			<b>2</b>	<b>2</b>	<b>62</b>	<b>2</b>	<b>14</b>	<b>101</b>	<b>6</b>	<b>189</b>	<b>75</b>	<b>264</b>	<b>3</b>
Phase I	1	8003	2	0	0	0	0	0	5	7	15	24	4
	1	8004	0	0	0	0	0	0	0	0	3	3	2
	1	8005	1	0	0	0	0	0	1	2	6	14	4
	1	8006	1	0	0	0	0	0	1	2	6	8	1
	1	8007	0	0	0	0	0	2	2	4	20	24	4
	1	8009	0	0	0	0	0	3	18	21	24	47	1
	2	8204	1	0	0	0	0	0	0	1	2	3	0
	2	8205	0	0	0	0	0	0	0	0	3	5	2
	2	8206	0	0	2	0	0	1	1	4	8	12	0
	2	8207	0	0	0	0	0	1	1	2	6	8	0
	2	8209	1	0	0	0	0	0	0	1	10	11	0
	2	8210	0	0	0	0	0	1	0	1	7	9	0
	2	8210/8211	0	0	0	0	0	0	1	1	2	3	0
	2	8211	0	0	0	0	0	0	1	1	7	8	0
	2	8212	0	0	0	0	0	2	2	2	5	7	0
	2	8215	0	0	2	0	0	1	1	4	3	7	0
	3	8303	6	0	0	0	0	0	1	7	11	18	1
	3	8304	11	0	0	0	0	1	0	12	11	23	3
	3	8305	0	0	0	0	0	0	1	1	8	9	2
	3	8306	2	2	0	0	0	0	0	4	8	12	0
	3	8307	4	1	0	0	0	0	2	7	28	35	12
	3	8308	6	3	0	0	0	3	6	18	87	96	11
<b>Totals</b>			<b>35</b>	<b>6</b>	<b>4</b>	<b>0</b>	<b>0</b>	<b>13</b>	<b>44</b>	<b>102</b>	<b>280</b>	<b>386</b>	<b>47</b>
Phase II	1	8010	0	0	0	1	0	0	1	2	14	18	3
	1	8011	0	0	0	0	0	0	0	0	1	2	0
	1	8019	0	0	0	0	0	0	0	0	2	2	0
	1	8020	1	0	0	0	2	0	4	7	16	23	5
	1	8015	0	0	0	0	0	1	6	7	37	44	4
	1	8016	0	0	0	0	0	1	6	7	45	52	13
	1	8017	0	1	0	0	1	0	2	4	29	35	12
	3	8309	0	0	0	0	0	0	1	1	10	11	0
	3	8310	3	0	0	0	0	1	1	5	13	18	7
	3	8311	0	0	0	0	0	0	0	0	3	3	0
	3	8312	1	0	0	0	0	1	2	4	4	8	2
	3	8313	0	0	0	0	0	0	0	0	2	2	0
	3	8315	0	0	0	0	0	1	1	2	7	9	0
	3	8316	1	0	0	0	0	0	2	3	6	9	2
	3	8317	0	0	0	0	0	2	1	3	16	19	6
	3	8318	0	0	0	0	0	0	1	1	1	1	0
	3	8319	2	0	1	0	0	0	0	3	2	5	2
	3	8321	0	0	0	0	0	0	6	6	52	58	18
	3	8322	0	0	0	0	0	0	2	2	22	24	2
	3	8323	0	0	0	0	0	0	0	0	2	2	1
<b>Totals</b>			<b>8</b>	<b>1</b>	<b>1</b>	<b>1</b>	<b>3</b>	<b>7</b>	<b>36</b>	<b>57</b>	<b>284</b>	<b>345</b>	<b>77</b>
Phase III	1	8018/8100	0	0	0	0	0	0	1	1	1	2	0
	1	8101	0	0	0	0	0	0	0	0	1	1	0
	1	8102	0	0	0	0	0	0	0	0	0	0	0
	1	8103	0	0	0	0	1	1	0	2	1	3	0
	1	8104 (=8109)	0	0	0	0	0	0	1	1	14	15	4
	1	8105	0	0	0	0	0	0	0	0	4	4	1
	1	8110	0	0	0	0	0	0	1	1	9	10	0
	1	8111	0	0	0	0	0	0	1	1	1	1	0
<b>Totals</b>			<b>0</b>	<b>0</b>	<b>0</b>	<b>0</b>	<b>1</b>	<b>1</b>	<b>4</b>	<b>6</b>	<b>31</b>	<b>36</b>	<b>5</b>

<< Table 5, vertebrate fauna remains from Hang Trong (2009 & 2010) analysed from trenches 1-3 and from all four identified phases. >>

Context	Phase	Distance above sea-level	Depth below surface	<i>Cyclophorus theodori</i>	<i>Cyclophorus unicus</i>	<i>Cyclophorus</i> cf. <i>cambodjensis</i>	<i>Cyclophoru</i> s spp.	<i>Cryptozona</i> cf. <i>chrysoraphe</i>	<i>Zonitidae</i> cf. <i>Oxychilus</i> sp.	<i>Ellobium aurisjudae</i>	<i>Camaena</i> sp.	<i>Unio</i> spp.	<i>Cerithrium</i> sp.	<i>Amphidromus</i> sp.	<i>Planorbis</i> sp.	Sinistral Gastropod (indet)	Gastropod A. (indet)	Gastropod B. (indet)	Total (unburnt) MNI count	Total (burnt) MNI count
8000 + 8001	Mixed	142.30-142.24	0-6	0	1	0	96	5	2	0	0	0	0	0	1	1	1	0	107	0
Total																			107	0
8003a	I	142.24-142.21	6-9	15	4	5	137	5	3	0	0	0	0	0	0	0	0	0	169	0
8003b		142.21-142.10	9-20	35	24	7	267	6	5	0	0	0	0	0	0	0	0	0	344	0
8004		142.10-142.08	20-22	7	6	2	69	0	0	0	0	0	0	0	0	0	0	0	84	0
8005		142.08-141.96	22-34	31	8	2	186	2	2	0	0	0	0	0	0	0	0	0	231	0
8006		141.96-141.81	34-49	43	8	4	187	2	2	0	0	0	0	0	0	0	0	0	246	155
8007a		141.81-141.75	49-55	38	24	7	69	0	1	0	1	1	0	0	0	0	0	0	141	0
8007b		141.75-141.68	55-62	5	5	0	110	1	1	0	0	0	1	0	0	0	0	0	123	66
8007c + 8009	141.68-141.63	62-67	8	4	2	32	0	0	0	0	0	0	0	0	0	0	0	46	3	
Total																			1384	224
8010	II	141.63-141.56	67-74	33	19	0	120	0	0	0	0	0	0	0	0	0	0	0	172	0
8011		141.56-141.47	74-83	26	27	4	101	0	0	0	0	0	0	0	0	0	0	0	158	4
8019 (S2)		141.47-141.33	83-97	55	28	7	276	0	0	0	0	1	0	1	0	0	0	0	368	8
8019 (S3)		141.33-141.20	97-110	34	43	13	321	0	0	1	0	1	0	2	0	0	0	1	416	10
8019 (S4)		141.20-141.08	110-122	27	8	10	452	0	0	0	0	1	0	1	0	1	0	0	500	0
8019 (S4)		141.08-140.97	122-133	38	60	15	229	0	0	0	0	0	0	0	0	0	0	0	342	4
Total																			1956	26
Totals				395	269	78	2652	21	16	1	1	4	1	4	1	2	1	1	3447	250

<< Table 6, Hang Trống mollusc sample column from Trench 1, north-facing section: taxonomic identification and quantification by depth and sample number. >>

Trench/Square	Context	Phase	Techno-typological analysis	Attribute analysis	Typological ID	Raw material	Length	Width	Thickness	Comments
Surface find	-		Secondary pebble fragment	-	Short axe	Shale	4.9	5.7	1.5	Transverse snap forms butt of tool
Deep Sounding	8101		Secondary flake with cortical platform	Complete flake	-	Limestone	2.7	2.4	0.8	-
Deep Sounding	8101		Secondary flake with plain platform	Flake fragment	-	Limestone	2	1.7	0.5	Stepped
Deep Sounding	8101		Burnt pebble fragment	?	-	Limestone	4.7	2.7	1.4	Burnt (natural?)
Deep Sounding	8101		Secondary flake with crushed platform	Complete flake	-	Limestone	4.3	8.1	1.7	-
Deep Sounding	8101		Flake core with single platform	Core	-	Limestone	6	6.2	3.2	4 removals
Deep Sounding	8103		Secondary pebble fragment	Pebble tool	Short axe	Limestone	3.8	6.1	1.6	Transverse snap forms butt of tool
Deep Sounding	8104		Pebble half fragment	Core/Pebble tool	-	Sandstone	3.1	3.7	1.4	Transverse snap
Deep Sounding	8104	III	Pebble fragment burnt	Complete flake	-	Limestone	4.1	2.5	1.4	burnt
Deep Sounding	8104		Secondary bifacial pebble tool	Pebble tool	Short axe	Shale	2.7	5.3	1.5	Transverse snap forms butt of tool; bifacial edge trimming
607/709	Cleaning		Secondary fragment	-	Scraper	Limestone	6.2	7.3	1.8	Retouched natural fragment; transverse snap
606/709	8109		Tertiary silet flake with crushed platform	Complete flake	-	Limestone	3.4	2.5	1.9	Half-moon snaps
607/709	8110		Secondary pebble fragment	Pebble tool	Short axe	Shale	3.7	4.9	1.3	Flattened pebble
606/709	8110		Plain platform tertiary flake	Complete flake	-	Shale	3.2	3.4	0.9	-
607/709	8110		Secondary flake with plain platform	Complete flake	-	Shale	1.7	3.1	0.6	Off pebble
606/709	8110		Secondary flake with cortical platform	Complete flake	-	Limestone	2.2	2.9	0.7	Off pebble
609/709	8110/8105		Tertiary flake with crushed platform	Retouched flake	Scraper	Quartzite	2.2	2.5	0.7	Stepped
601/712	8200		Flaked pebble fragment	Core	Core?	Limestone	3.8	5.3	6.3	Steep edged; platform on each end; stepped flaking removals
601/712	8200		Distal tertiary flake fragment	Complete flake	-	Shale	2	2.4	0.4	Transverse snap
601/712	8200		Distal tertiary flake fragment	Complete flake	-	Shale	1.6	2.5	0.4	-
602/712	8201	Mixed	Tertiary flake with plain platform	Complete flake	-	Limestone	1.7	1.3	0.6	Edge snaps
603/712	8200/8201		Tertiary flake with crushed platform	Complete flake	-	Limestone	3.5	2.7	0.6	Oblique transverse snap; half-moon snaps
603/712	8200/8201		Tertiary flake with crushed platform	Complete flake	-	Limestone	2.9	2.6	0.6	Edge snaps
601/712	8203		Distal primary flake fragment	Flake fragment	-	Shale	3.2	3.2	0.8	Transverse snap
602/712	8203		Elongate pebble fragment	-	-	Shale	5.8	4.9	2.6	Natural?
603/712	8207		Tertiary shatter fragment	-	-	Shale	2.4	1.5	0.7	Natural?
603/712	8209	I?	Tertiary flake with plain platform	-	-	Shale	2.5	4.5	0.8	-
603/712	8211		Distal tertiary flake/blade fragment	Complete flake	-	Shale	2	1.3	0.4	Transverse snap
609/714	8300	Mixed	Secondary pebble flaked fragment	Pebble tool	Short axe	Limestone	7.8	9.5	3.2	Transverse snap forms butt of tool
610/715	8303		Secondary flake with crushed platform	Complete flake	-	Limestone	2.6	2.7	0.3	Edge snaps
610/715	8303		Secondary pebble fragment	-	-	?	4.1	3	1.2	Heavily burnt
610/715	8303		Secondary flake with crushed platform	Complete flake	-	Limestone	1.4	1.6	0.1	Edge snaps
610/715	8304		Distal tertiary flake fragment	Flake fragment	-	Shale	2.2	1.8	0.3	Transverse snap and edge flaking off it
610/715	8304		Tertiary shatter fragment	-	-	Limestone	1.7	2.4	0.5	Natural?
610/714	8304		Secondary fragment	-	-	Shale	-	-	-	Natural?
610/714	8304		Flake with plain platform	Complete flake	-	Limestone	3.4	4.4	0.7	Stepped; edge damage
610/714	8304		Tertiary pebble fragment	-	-	Limestone	-	-	-	Natural?
610/715	8305		Pebble single platform flake core	Core	Core	Limestone	3.5	5.3	9.4	Elongate pebble; 4 removals
610/715	8305		Tertiary flake with crushed platform	Complete flake	-	Shale	1.6	1.3	0.3	-
610/715	8306		Distal tertiary flake fragment	?	-	Shale	1	1.8	0.2	Transverse and edge snaps
610/714	8307	I	Secondary flake with cortical platform	Complete flake	-	Shale	3.6	2.6	0.8	-
610/714	8307		Distal primary blade fragment	Complete flake	Burin	Shale	4.2	1.5	0.8	-
610/714	8307		Tertiary flake fragment with cortical	Flake fragment	-	Shale	1.5	1.2	0.4	-
610/714	8307		Distal tertiary flake fragment	Flake fragment	-	Shale	1.3	1.2	0.4	-
610/714	8307		Secondary shatter fragment	-	-	Shale	3.2	1.8	0.6	Natural?
609/716	8308		Retouched pebble fragment	Pebble tool	Short axe	Shale	6.4	6.4	1.8	Flattish pebble
609/716	8308		Tertiary flake fragment with cortical	-	-	Shale	6.5	4.4	1.1	Half-moon snaps
610/715	8308		Flaked secondary pebble chopper/core	Pebble tool	Chopper	Shale	6.4	9.8	3.9	6 flake removals if this was core
610/715	8308		Secondary flake with crushed platform	Complete flake	-	Shale	2.3	2.8	0.5	-
610/715	8308		Secondary flake with crushed platform	Complete flake	-	Shale	1.7	2.8	0.3	-
610/715	8308		Secondary flake with plain platform	Complete flake	-	Shale	2.2	3.4	0.8	-
610/715	8308		Pebble	Pebble tool	Pounder	Limestone	9.6	9.1	6.7	Battering + flaking damage (opposed ends)
610/714	8309		Tertiary flake with plain platform	Complete flake	-	Shale	1.3	1.8	0.5	-
610/715	8312		Tertiary flake with crushed platform	Complete flake	-	Shale	4.1	3.6	0.9	Lateral snaps
610/715	8312		Primary flake with cortical platform	Complete flake	-	Shale	2.4	3.7	0.7	Half-moon snaps
610/715	8312		Tertiary blade/flake with crushed	Complete flake	-	Shale	2.5	1.4	0.3	-
610/715	8312		Tertiary flake with crushed platform	Flake fragment	-	Shale	2.1	4.1	0.5	Large half- moon snaps
610/715	8312		Tertiary potlid	-	-	Shale	1.3	2	0.3	Natural?
610/715	8313		Tertiary flake with crushed platform	Complete flake	-	Shale	2.3	4.5	0.8	-
610/714	8315		Secondary silet flake with cortical	Complete flake	-	Shale	3.4	2.6	1.1	-
610/714	8315		Secondary flake with cortical platform	Complete flake	-	Shale	3.6	6.3	1.4	-
610/714	8315		Tertiary flake with crushed platform	Complete flake	-	Shale	3.2	2.5	0.7	-
610/714	8315		Secondary shatter fragment	-	-	Shale	2.8	2.1	0.6	Natural?
610/714	8315		Tertiary flake fragment	Flake fragment	-	Limestone	1.9	3.6	0.6	Snapped around whole edge
610/714	8315		Distal tertiary flake fragment	Flake fragment	-	Limestone	2.3	2.3	0.3	Transverse and edge snaps
610/714	8315		Tertiary potlid	-	-	Limestone	1.4	1.7	0.3	Edge snaps (natural)
610/714	8315		Tertiary flake with plain platform	-	-	Shale	1.3	2.5	0.5	Stepped
610/714	8315		Tertiary shatter fragment	-	-	Shale	1.6	2.1	0.3	Flaking debris or natural?
610/714	8315		Distal tertiary flake fragment	Flake fragment	-	Shale	1.2	2.1	0.3	Transverse snap
610/714	8316	II	Tertiary flake with crushed platform	Complete flake	-	Shale	2.2	2.2	0.2	Hinged
610/714	8317		Primary silet flake with crushed platform	Flake fragment	-	Shale	2.4	1.5	0.7	Half-moon snaps
609/715	8322		Secondary flake with plain platform	Flake fragment	-	Shale	0.9	3.2	0.8	-
609/715	8322		Tertiary shatter fragment	-	-	Shale	1.8	2.5	0.2	Natural?
609/715	8322		Tertiary shatter fragment	-	-	Shale	2.6	1.7	0.5	Natural?
609/715	8322		Tertiary shatter fragment	-	-	Shale	3	1.8	0.8	Natural?
609/715	8322		Tertiary flake with plain platform	-	-	Shale	2.8	2.9	0.7	-
609/715	8322		Tertiary flake with crushed platform	-	-	Shale	2.2	2.6	0.9	-
609/715	8322		Tertiary flake with crushed platform	-	-	Shale	2.4	3.2	0.5	-
609/715	8322		Distal tertiary flake fragment	-	-	Shale	1.8	1.8	0.6	-
609/715	8322		Tertiary flake with plain platform	Flake fragment	-	Shale	1.4	1.4	0.3	-
609/715	8322		Distal tertiary flake fragment	Complete flake	-	Shale	1.2	2	0.5	Transverse snap
609/715	8322		Distal tertiary flake fragment	-	-	Shale	1.4	2.4	0.3	Transverse snap
609/715	8322		Distal tertiary flake fragment	Flake fragment	-	Shale	1.4	1.5	0.2	Transverse snap
609/715	8322		Tertiary flake with plain platform	-	-	Shale	1.6	2.2	0.4	-
609/715	8322		Distal tertiary flake fragment	Flake fragment	-	Shale	1.2	1.7	0.4	-
609/715	8322		Distal tertiary flake fragment	-	-	Shale	1.3	1.5	0.2	-
610/715	8322		Secondary pebble fragment, retouched	Pebble tool	Short axe	Shale	4.6	6.4	1.8	Transverse snap forms butt of tool

<< Table 7, classification of material excavated in 2010 from Trench 1 (Phase III) and Trenches 2 and 3 (all represented deposition phases) using data generated separately by techno-typological and attribute analysis methods; classic typological referents are also given where relevant. >>